A brightening of Jupiter's auroral 7.80- μ m CH₄ emission during a solar-wind compression

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Enhanced mid-infrared emission from CH₄ and other stratospheric hydrocarbons have been observed coincident with Jupiter's ultraviolet auroral emission¹⁻³, which demonstrates that 2 auroral processes and the neutral stratosphere of Jupiter are coupled. However, the exact 3 nature of this coupling has remained an open question. Here, we present a time series of 4 Subaru-COMICS images of Jupiter at 7.80 µm measured between January 11-14th, Febru-5 ary 4-5th and May 17-20th (UT) 2017, which show both the morphology and magnitude of 6 the auroral CH_4 emission to vary on daily timescales in relation to external solar-wind condi-7 tions. The southern auroral CH₄ emission increased in brightness temperature (T_b) by ΔT_b 8 = 3.8 ± 0.9 K between January 11th 15:50 UT - 12th 12:57 UT during a predicted solar-9 wind compression. During the same compression, the northern auroral emission exhibited 10 a dusk-side brightening, which mimicks the morphology observed in the ultraviolet auro-11 ral emission during periods of enhanced solar-wind pressures^{4,5}. These results suggest that 12 changes in external solar wind conditions perturb the jovian magnetosphere such that ener-13

¹⁴ getic particles are accelerated into the planet's atmosphere, deposit their energy as deep as ¹⁵ the neutral stratosphere and modify the thermal structure, abundance of CH_4 or the pop-¹⁶ ulation of energy states of CH_4 . We also find that the northern and southern auroral CH_4 ¹⁷ emission evolved independently between the January, February and May images, as has been ¹⁸ observed at X-ray wavelengths over shorter timescales⁶ and at mid-infrared wavelengths over ¹⁹ longer timescales⁷.

7.80-µm Subaru-COMICS (Cooled Mid-Infrared Camera and Spectrograph) images were 20 obtained from January 11-14th, February 4-5th and May 17-20th 2017 (UT). A subset of images 21 recorded are shown in Figures 1 and 2, which respectively show southern and northern polar pro-22 jections at times when the southern auroral region (henceforth 'SAR', between 330-60°W in Sys-23 tem III) and northern auroral region (henceforth 'NAR', centred at 180°W in system III longitude) 24 were visible on the disk of Jupiter. These images demonstrate variability of both the magnitude 25 and morphology of the 7.80- μ m CH₄ emission over timescales of days to months. Further details 26 of the measurements and processing are provided in the Methods and Supplementary sections. 27

In terms of the morphology, the strongest 7.80- μ m emission in both auroral regions appears enclosed inside the statistical mean of the ultraviolet emissions of the main oval⁸. Figure 3 shows the results of ionosphere-to-magnetosphere mapping model calculations (see Methods) and demonstrates that the positions of strongest CH₄ emission in the auroral regions predominantly correspond to radial distances of >95 R_J (beyond the dayside magnetopause⁹ and potentially on open field lines). The exception is the morphology of the emission in the NAR at 16:13 UT on January

12th (Figure 2a), when a poleward, duskside feature of stronger emission parallel to the eastern 34 boundary of the statistical oval was observed. This feature was not present less than 24 hours later 35 (Figure 2b) and we have ruled out variable atmospheric seeing conditions between these two nights 36 as the source of this intermittent morphology (see Supplementary Figure 2). A similar morphology 37 of the ultraviolet auroral emission, described as the 'duskside active region', has also been observed 38 during periods of enhanced solar-wind pressures and has been attributed to duskside/nightside re-39 connection associated with the Vasyliunas or Dungey cycles or velocity shears caused by changing 40 flows on the nightside magnetospheric flank^{4,5,10}. Indeed, ionosphere-magnetosphere mapping 41 calculations map 73°N, 155°W (an example location covered by the duskside feature) to $\sim 100R_J$ 42 at a local time of 19.0 hr. Unlike the NAR, the SAR does not appear to exhibit any smaller-scale 43 morphology although its position at a comparably higher latitude than the NAR does reduce the 44 effective spatial resolution and the ability to resolve smaller-scale features. In contrast to previous 45 studies^{7,11}, we find no obvious movement in the longitudinal position of the southern auroral CH_4 46 emission in the images presented in this work. 47

In order to quantify temporal changes in the magnitude of the auroral emission and its relation to solar-wind conditions, we calculated the residual radiance between each auroral region and a lower-latitude zonal-mean, henceforth named the auroral-quiescent residual (see Methods). Figure 4 compares the auroral-quiescent residual and uncertainty for both auroral regions and the results of a solar-wind propagation model (see Methods). The solar-wind propagation model predicts the arrival of a solar-wind compression at Jupiter at approximately 22:00 UT on January 11th, when the dynamical pressure was predicted to have increased from <0.1 nPa to 0.7 nPa. The auroral-

quiescent residual increased from $T_b = 8.0 \pm 0.3$ K on January 11th 15:50 UT to 11.8 \pm 0.5 K 55 on January 12th 12:57 UT: a net increase of 3.8 \pm 0.6 K in T_b or a ~25% increase in radiance. 56 While the viewing geometries of the SAR differ between these two respective images, forward-57 model calculations of the 7.80-µm emission (see Methods) at these two geometries differ by only 58 0.7 K in T_b and thus cannot explain all of the observed change. From January 12th 12:57 UT to 59 January 14th 12:33 UT, the SAR returned to a similar brightness on January 14th as was observed 60 pre-compression and was a similar brightness in all subsequent measurements (although variability 61 intermediate of these measurements cannot be ruled out). 62

The NAR was not visible on the disk of Jupiter in the images measured on January 11th 63 (before the solar-wind compression) and so we do not know whether it also brightened during the 64 same solar-wind compression. However, the aforementioned duskside-active emission captured by 65 COMICS on January 12th 16:13 UT (Figure 2a) occurred shortly after the solar-wind compression, 66 which reiterates that this morphology is likely driven by enhanced solar-wind conditions and their 67 perturbing effect on the nightside magnetosphere. From January 12th 16:13 UT to January 13th 68 12:30 UT, the auroral-quiescent residual of the NAR was constant in time within uncertainty and 69 subsequently decreased significantly to 1.2 ± 1.1 K. Similarly, measurements in May show the 70 NAR emission to be weak and comparable with, if not weaker, than lower-latitude regions. From 71 May 18th to May 19th, there was a marginal increase in the emission in the NAR during a small 72 $(\Delta p_{dyn} \sim 0.2 \text{ nPa})$ solar-wind compression however, the change in emission was not significant 73 with respect to measurement uncertainty. Without measurements between the dates of January 74 13th, February 5th and May 18th, it is uncertain whether the NAR emission was consistently 75

weaker in time or whether it exhibited short-term (daily/weekly) variability and the measurements 76 by chance captured periods of weaker emission. However, we favour the latter possibility given that 77 the February-5th and May-17th measurements were preceded by \geq 7 days of steady, low pressure 78 $(p_{dyn} < 0.05 \text{ nPa})$ solar-wind conditions. We note the results of a recent study by Kita et al. 79 2016¹⁵, which showed the total auroral power during a solar-wind compression exhibited a positive 80 correlation with the duration of steady, quiescent solar-wind conditions preceding the compression. 81 We also note that the northern auroral C_2H_6 emission weakened during periods of low solar activity 82 in previous work, which similarly suggests a connection with solar-wind conditions on longer 83 timescales¹³. 84

The daily variability of the southern auroral CH₄ emission is suggestive the source of the 85 variability occurs in the upper stratosphere/mesophere to thermosphere region (10 - 1 μ bar), where 86 the thermal inertial timescales are much shorter (~ 4 weeks at 1 µbar¹⁴) compared to the lower 87 stratosphere (\sim 30 weeks at 1 mbar¹⁴). We suggest the observed changes in CH₄ emission result 88 from either: 1) variable auroral-related heating of the 10- to 1-µbar level, 2) auroral-driven changes 89 in the vertical profile of CH₄ near its homopause at $\sim 1 \mu bar$, 3) variable non-LTE effects that 90 modify the population of energy states of CH_4 or 4) some combination of 1-3. In order to explore 91 the first two possibilities and determine what magnitude and type of changes in the vertical profiles 92 of temperature or CH₄ could yield a 7.80- μ m Δ T_b = 3 - 4 K increase, we performed a series of 93 radiative-transfer calculations using NEMESIS (see Methods). 94

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As shown in Supplementary Figures 5a-b, assuming the CH₄ abundance is held fixed, a 3

- 4 K change in T_b would require either: 1) the pressure level of the mesosphere-thermosphere 96 transition to move deeper in the atmosphere by approximately a pressure-scale height or 2) the 97 lapse rate in the thermosphere to increase by a factor of 2. The former corresponds to a total, atmo-98 spheric temperature increase of more than 100 K at the 0.5-µbar level, assuming a thermospheric 99 lapse rate similar to that measured during Galileo's descent¹⁵, whereas the latter corresponds to 100 a total, atmospheric temperature increase of ~ 20 K at 0.5 µbar. In steady state, thermospheric 101 general circulation models show that the mesosphere-thermosphere transition pressure is deeper 102 in the auroral regions compared to non-auroral regions^{16,17}. Yates et al., 2014¹⁸ performed time-103 dependent thermospheric circulation modelling to investigate the response of the thermospheric 104 structure and circulation to solar-wind compressions and expansion events. Between steady and 105 compressed solar-wind conditions, the model predicted a ~ 20 -K warming and increase in lapse 106 rate at $\sim 70^{\circ}$ N due to increased rates of joule heating at pressures lower than 1 µbar (with the lower 107 model boundary set at 2 μ bar). This is consistent with the two-fold increase in thermospheric lapse 108 rate required to brighten the 7.80- μ m emission by $\Delta T_b = 3 - 4$ K, as detailed above. 109

As shown in Supplementary Figure 5c, assuming a fixed vertical temperature profile, increasing the altitude of the CH₄ homopause (with respect to the Moses et al., 2005²⁰ Model A CH₄ profile) by greater than a pressure-scale height would yield a 3 - 4 K increase in T_b at 7.80-µm. At the 0.2-µbar level, this would correspond to a volume mixing ratio (VMR) increase of ~10⁻⁴. In solving the vertical continuity equation assuming the change in VMR is driven entirely by advection and not a chemical source (i.e. $w = (-\Delta X/\Delta t)/(\Delta X/\Delta z)$, where w is the vertical velocity, X is the VMR, t is time and z is height), a change in vertical wind of 2.7 cm s⁻¹ with respect to the steady state would be required. The Bougher et al., 2005 thermospheric model ¹⁶ predicts vertical winds at \sim 70°S of approximately 50 cm s⁻¹ at the 0.2-µbar level in steady state and thus a change in vertical wind of 2.7 cm s⁻¹ is reasonable. A higher-altitude homopause of CH₄ (and other hydrocarbons) in Jupiter's auroral regions was also found to optimize the consistency between Juno and Hisaki measurements²⁰.

Non-LTE effects are also likely important at the altitudes where the source of variability 122 has been inferred and/or could itself be the driver of the observed variability. In the absence of a 123 strong radiation source, 'classical' non-LTE effects become non-negligible at pressures lower than 124 0.1 mbar, where collisional timescales become longer than the spontaneous radiative lifetime $^{21-23}$. 125 Without a sufficient number of thermal collisions, the population of rotational and vibrational 126 energies deviates from the translational energy population and thus can no longer be described 127 as a Boltzmann distribution. In comparison to non-auroral regions, the upper-stratospheric heating 128 present in Jupiter's auroral regions^{7,19,25} also yields a larger contribution of photons at mid-infrared 129 wavelengths from pressure levels where classical non-LTE processes become non-negligible. In 130 addition, currents of electrons and ions in Jupiter's auroral regions and the resulting charged-13 particle collisions and dissociative recombinations may induce a non-Boltzmann population of 132 the excited energy states of CH_4 . A further process might be ' H_3^+ -shine', where the downward 133 flux of H_3^+ emission in lines in the 3- to 4-µm range 'pump' overlapping CH₄ ν_3 lines, exciting 134 the vibrational modes and thereby modifying the population of lines responsible for the v_4 band 135 at ${\sim}7.80~\mu m^{26}.$ Modeling of the aforementioned non-LTE processes will be the subject of future 136 work. 137

We cannot distinguish between temperature, CH_4 abundance, or non-LTE effects in driv-138 ing the variable CH₄ emission observed between January 11-12th 2017. Nevertheless, either of 139 these processes describes a direct coupling of the neutral stratosphere in Jupiter's auroral regions 140 to the external magnetosphere of Jupiter and solar-wind environment. While daily variability of 141 the northern auroral C_2H_4 and C_2H_6 emission has been observed in previous studies^{27,28}, we be-142 lieve the results presented in this work represent a significant advance in understanding of this 143 phenomenon. Firstly, the availability of solar-wind measurements and their modelled propagation 144 to Jupiter's orbit allowed the variability of the CH₄ emission to be tentatively linked to exter-145 nal solar-wind changes and their perturbing effect on the magnetosphere. Secondly, COMICS 146 imaging at high-diffraction limited spatial resolutions have allowed the morphology of the CH₄ 147 emission and its variability to be resolved at finer spatial details and mapped to the outer magneto-148 sphere/magnetopause using ionosphere-to-magnetosphere mapping calculations. Auroral-related 149 heating and chemistry dominate the forcing of the thermal structure and composition at Jupiter's 150 poles 7, 19, 25 and the results of this paper suggest these processes are directly connected to the exter-151 nal magnetosphere. This phenomenon therefore could be ubiquitous for rapidly-rotating Jupiter-152 like exoplanets with an internal-plasma source around a magnetically-active star²⁹. In particular, 153 MHD (magnetohydrodynamic) simulations of a hot-Jupiter at close orbital separations of 0.05 AU 154 from its host star predict auroral powers of several orders of magnitude larger than on Earth and to 155 affect both polar and equatorial regions³⁰. The coupling of the neutral stratosphere and magneto-156 sphere of Jupiter presented in this work may therefore be a process of importance in the near-future 157 characterization of Jupiter-like exoplanets from the James Webb Space Telescope and/or directly-158

¹⁵⁹ imaged planets whose atmospheres are predominantly sensed at higher latitudes.

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Author contributions JAS led the analysis of the observations and the preparation of this letter. Co-244 authors GSO and YK were principal investigators of the awarded telescope time from which the presented 245 data was measured. JAS and co-authors GSO, YK, TMS and TF participated in the measurements at the 246 Subaru telescope. Co-author JF performed the reduction and calibration of the images. Co-authors CT and 247 MFV provided model output, which was invaluable in the interpretation of the results. Co-author PGJI is the 248 lead developer of the NEMESIS code, which was adopted in the paper for radiative transfer forward-model 249 calculations. All remaining authors contributed to the interpretation of the results and the preparation of this 250 publication. 251

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COMICS 7.8-µm images The COMICS (the COoled Mid-Infrared Camera and Spectrograph^{1,2}) 256 instrument is mounted at the Cassegrain focus of the Subaru Telescope, which is located at the 257 Mauna Kea Observatory (approximately 4.2 km above sea level). Subaru's 8.2-metre primary 258 aperture provides a diffraction-limited spatial resolution of ~ 0.24 " at 7.8 μ m, which corresponds 259 to a latitude-longitude footprint of approximately 2.5° x 2° at $\pm 70^{\circ}$ in latitude. COMICS provides 260 both imaging and spectroscopic capabilities over a spectral range of approximately 7 to 25 μ m. 261 Images are measured on a 320 x 240 array of Si:As BIB (blocked impurity band) pixels each with 262 a scale of 0.13", which provides a total field-of-view (FOV) of 42" x 32". Images can be measured 263 over a number of discrete filters in both the N band (7 to 13 μ m) and Q band (17 to 25 μ m). In this 264 work, we focus on images obtained in the 7.8- μ m filter, which is sensitive to Jupiter's stratospheric 265 CH_4 emission (Supplementary Figure 3). Images were measured on January 11-14, February 4-266 5 and May 17-20 2017 (UTC). Measurements were performed during periods when Jupiter was 267 available at airmasses lower than 3. The full disk of Jupiter (with equatorial diameters of \sim 36" 268 in January, ~ 39 " in February and ~ 42 " in May) could not be measured in a single image by the 269 COMICS field-of-view (FOV). In January and February measurements, the full disk of Jupiter was 270 measured using a 2 x 1 mosaic of individual images centred at Jupiter's mid-northern and mid-27 southern latitudes. In May, a 2 x 2 mosaic was conducted due to Jupiter's larger size during this 272 time period. For each individual image, A-frames (of Jupiter) and B-frames (dark sky 60" north 273 of Jupiter) were continuously recorded over a total exposure time of 20 seconds. Further details of 274 the measurements presented in this work are provided in Supplementary Table 1. 275

Imaging processing, calibration and error handling Images were processed and calibrated us-276 ing the Data Reduction Manager (DRM). A-B subtraction was performed to remove telluric sky 277 emission. The resulting images were then divided by a 'bad pixel mask' accounting for corrupted 278 pixels (due to cosmic ray damage, bright star saturation, manufacturer flaws etc.) and a flatfield in 279 order to remove variations in pixel-to-pixel sensitivity across the detector. A limb-fitting procedure 280 was used to assign latitudes, longitudes and local zenith angles to each pixel on the disk of Jupiter, 281 using the known sub-observer latitude and longitudes at the time of each exposure. The abso-282 lute radiometric calibration of the images and correction for telluric absorption was conducted by 283 scaling the observed lower-latitude zonal-mean brightness to those measured by Cassini's CIRS³ 284 instrument during the 2001 flyby. This procedure is described in greater detail in Fletcher et al., 285 2009^4 . We chose this method of calibration since experience with past mid-infrared images of 286 Jupiter and Saturn has demonstrated that the radiometric calibration using a standard star pro-287 vides inconsistency between datasets obtained on different nights^{4,5}. As detailed further in the 288 'Auroral-quiescent residual calculations' section of Methods, our analysis of the images involved 289 comparing the relative brightness of the auroral regions with a lower-latitude region over time, 290 which negates errors introduced by offsets in the absolute calibration between nights. The reduced 291 and radiometrically-calibrated images are shown in Supplementary Figure 1 in units of brightness 292 temperature (T_b) at 7.80 µm. Portions of the image within 6 pixels (or approximately 0.8") of 293 the assigned limb were removed as a conservative means of removing the effects of seeing and 294 diffraction in blurring dark sky together with emission from Jupiter. The noise-equivalent spec-295 tral radiance (NESR) was calculated by finding the standard deviation emission of dark-sky pixels 296

more than 1.5" (or approximately 12 pixels) away from the planet. This was calculated for each 297 image to capture changes in sensitivity due to variations in airmass and telluric atmospheric con-298 ditions between measurements. A centre-to-limb variation correction in the longitudinal direction 299 was applied to correct for the foreshortening and limb-brightening such that longitudes at different 300 viewing geometries on different nights could be more readily compared. A power-law fit, of the 301 form $\log R = a \log \mu + b$, where R is radiance, $\mu = \cos \theta$ and θ is the zenith emission angle, was 302 performed in each latitude band in order to derive a centre-to-limb correction factor. For January 303 and February measurements, we performed the power-law correction using the January 11th 15:50 304 UT image (Supplementary Figure 1a) in the northern hemisphere and the January 13th 12:30 UT 305 image (Supplementary Figure 1d) in the southern hemisphere. For May measurements, the May 306 17th 09:40 UT and May 18 09:35 UT images (Supplementary Figures 1i, j) were similarly chosen 307 to perform the power-law correction in the northern and southern hemispheres, respectively. These 308 specific images were chosen since they best capture non-auroral longitudes in each hemisphere. 309

Ionosphere to magnetosphere mapping We adopted the magnetosphere-ionosphere mapping 310 calculation by Vogt et al.^{6,7} to map a location on the planet in planetocentric latitude and sys-311 tem III longitude to its position in radial distance and local time in the jovian magnetosphere. The 312 calculation is performed by imposing magnetic flux equivalence of a specified region at the equator 313 to the area at which it maps in the ionosphere assuming a given internal field model. For this work, 314 we adopted the VIPAL (Voyager Io Pioneer Anomaly Longitude) internal field model⁸ due to its 315 validity in both the northern and southern hemispheres and to larger ($\sim 95 \text{ R}_J$) radial distances. 316 Stepping through latitude and longitude in increments of 1° poleward of $\pm 45^{\circ}$ in latitude, the 317

ionosphere-to-magnetosphere mapping calculation was performed to derive the local time and distance within the magnetosphere at each location. Regions enclosed within the statistical ultraviolet oval for which the calculation did not produce a real value were interpreted as mapping beyond the 95 R_J limit of the model, which also marks the estimated position of the dayside magnetopause⁹. This calculation was used to derive the contours of distance shown in Figure 3.

Auroral-quiescent residual calculations Figure 3 demonstrates the areas denoted by 'Region A' 323 and 'Region L' at both high-northern and high-southern latitudes. Region A (for 'auroral') was 324 chosen as a sub-region of the auroral regions that mapped to the outer magnetosphere and was 325 commonly sampled by all measurements presented in Figures 1 and 2. Region L was chosen as a 326 lower latitude region away from the area of auroral influence, which is sampled at $\mu = cos(\theta_{emm})$ 327 (where θ_{emm} is the zenith emission angle on Jupiter) in the range 0.4 < μ < 1 in each image. By 328 calculating the residual between Region A and Region L, any inconsistencies in the radiometric 329 calibration from one night to the next are effectively removed, which would otherwise affect a 330 comparison of the absolute radiance in time. The mean radiances within Region A and Region L 331 were calculated. The 1- σ uncertainty on the mean radiance in each region was chosen to be the 332 larger of: 1) the NESR of each image (see Imaging processing, calibration and error handling) 333 scaled by $1/\sqrt{n_p}$ where n_p is the number of pixels averaged or 2) the standard deviation on the 334 mean radiance in the region. The radiances and uncertainties were then converted to brightness-335 temperature units and the brightness- temperature residual and uncertainty were calculated. 336

Solar-wind propagation model The Juno spacecraft continues to provide information on the
 magnetic and charged particle fields whilst performing 53.5-day orbits inside Jupiter's magne-

tosphere. However, the Juno spacecraft cannot provide *in-situ* measurements of the external solar-339 wind conditions outside Jupiter's magnetosphere. In the absence of such measurements, we look 340 instead to modelling results. A solar-wind propagation model¹⁰ was adopted to calculate the solar-341 wind dynamical pressure ($p_{dyn} = \rho v^2$, where ρ is the density and v is the velocity of the solar wind) 342 impinging on Jupiter's magnetosphere. This model is used extensively by the outer planets mag-343 netosphere community¹¹⁻¹³ in the absence of in-situ measurements of the solar-wind conditions. 344 The model adopts hourly measurements of the solar wind and magnetic field at Earth's bow-shock 345 nose from OMNI¹⁴ as input and then performs 1-D magnetohydrodynamic (MHD) calculations to 346 model the solar-wind flow out to Jupiter's bow-shock. The 1-D model prediction of a 3D problem 347 can introduce uncertainties on the arrival time and magnitude of dynamical pressure of solar wind 348 compressions. When the Earth-Sun-Jupiter angle is less than $\pm 50^{\circ}$, the uncertainty of the arrival 349 time of the solar wind shock is less than ± 20 h and that of the maximum dynamic pressure is 350 38%¹⁵. Given Earth-Sun-Jupiter angles were between 80 - 120° in the January-February 2017 time 351 range, we adopted a 48-hour time error on the solar wind propagation model results. In May 2017, 352 the Earth-Sun-Jupiter angle was approximately 18° and thus we assumed a time error of 20 hours 353 in the May time range. These values also appear commensurate with a statistical comparison of 354 1-D MHD predictions and solar-wind data measured by several spacecraft¹⁶. The aforementioned 355 error values are shown in Figure 4 to highlight the potential error to the reader. 356

Nemesis forward model calculations A single, broadband measurement of the CH_4 emission does not provide sufficient information to invert or *retrieve* atmospheric parameters and determine at what altitudes they vary. Nevertheless, we computed synthetic or 'forward-model' spectra for

a range of vertical profiles of temperature and CH₄ in order to explore what changes in those 360 atmospheric parameters could yield the observed, 7.80- μ m, $\Delta T_b = 3 - 4$ K brightening of the 361 southern-auroral region. The NEMESIS forward model and retrieval tool¹⁷ was adopted to com-362 pute forward-model spectra of the radiance in the COMICS 7.80-µm bandpass. Forward-model 363 spectra were computed using the line-by-line method using the sources of line information for 364 CH_4 , CH_3D and ${}^{13}CH_4$, C_2H_2 , C_2H_6 , NH_3 and PH_3 detailed in Table 4 of Fletcher et al., 2012¹⁸. 365 Calculations were performed using a square instrument function with a width of 0.04 cm⁻¹ (cho-366 sen based on a balance of a sufficiently high spectral resolution to resolve both weak and strong 367 emission lines whilst minimising computational expense) and subsequently convolved with the 368 COMICS 7.80-µm bandpass and the telluric transmission spectrum (see Supplementary Figure 369 2). The vertical profiles of temperature and CH_4 were varied as detailed below. The remaining 370 parameters of our model atmosphere, including the vertical profiles of C₂H₂, C₂H₄, C₂H₆, NH₃, 371 PH₃, were held constant since they have negligible effect on the spectrum in the 7.80-µm band-372 pass. Further details of the model atmosphere are provided in Sinclair et al., 2017¹⁹. It should be 373 noted that the current NEMESIS forward model assumes local thermodynamic equilibrium (LTE) 374 conditions, whereas, conditions in the auroral regions may have departed from LTE as discussed 375 in the main text. 376

³⁷⁷ Firstly, we kept the vertical profile of CH_4 and its isotopologues fixed to the 'model A' ver-³⁷⁸ tical profile from Moses et al., 2005²⁰. Starting from the temperature profile shown in Supplemen-³⁷⁹ tary Figure 4a, we modified the vertical temperature profile in the 0.1-mbar to 1-µbar range, which ³⁸⁰ includes the transition from the upper stratosphere/mesosphere to the thermosphere. The vertical

temperature gradient (or lapse rate) in thermosphere was fixed and the pressure level of the meso-38 sphere/thermosphere transition was varied as shown in Supplementary Figure 5a. For each profile, 382 a forward model was computed at the same viewing angle ($\mu = \cos(\theta_{emm}) = 0.205$, where θ_{emm} is 383 the emission angle) as Region A in the southern auroral region on January 12th 12:57 UT (during 384 the solar-wind compression). The synthetic spectrum was convolved with the 7.80-µm bandpass 385 (as detailed above) and converted into units of brightness temperature (T_b) . These T_b values are 386 shown in the legend in Supplementary Figure 5a. A further set of forward models and brightness 387 temperatures were similarly computed, where the pressure level of the mesosphere/thermosphere 388 transition was fixed at 0.2 µbar and the vertical temperature gradient (or lapse rate) was varied, as 389 shown in Supplementary Figure 5b. 390

³⁹¹ Secondly, we fixed the vertical profile of temperature as shown in Supplementary Figure 4a. ³⁹² Starting from the vertical profile of CH_4 derived from model A of Moses et al., 2005²⁰, the pressure ³⁹³ level of the methane homopause was varied as shown in Supplementary Figure 5c, a forward-³⁹⁴ model radiance in the 7.80-µm bandpass calculated and converted into brightness-temperature ³⁹⁵ units. These values are shown as the legend of Supplementary Figure 5c.

Data & Code availability The COMICS images presented in this work are now publically available on the SMOKA (Subaru Mitaka Okayama-Kiso Archive System, https://smoka.nao.ac.jp/) following an 18-month proprietary period since measurement. Reduced and calibrated images may be requested from author JAS with permission of the principal investigator of the awarded telescope time (see Acknowledgements). The Data Reduction Manager is a suite of IDL software designed for reduction and processing of planetary images and is available in compressed format from coauthor GSO upon request (glenn.s.orton@jpl.nasa.gov). The ionosphere-to-magnetosphere mapping calculation is also written in IDL and is available from co-author MFV (mvogt@bu.edu), upon request. Results of the solar wind propagation model in a specific time period may be requested from co-author CT (chihiro.tao@nict.go.jp), upon request. The NEMESIS forward model and retrieval tool is written in Fortran and is available as a GitHub repository: a user account for this repository may be requested from co-author PGJI (patrick.irwin@physics.ox.ac.uk).

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Southern-polar projections of Jupiter's 7.80- μ m CH₄ emission. Images were Figure 1 456 recorded by Subaru-COMICS on (a-c) January 11-14, (d) February 4 and (e-f) May 17, 457 20 2017. These are a subset of the observations shown in Supplementary Figure 1, 458 when the southern auroral region (330-60°W System III) was fully or partially visible on 459 the disk. Images are shown in brightness temperature units according to the bottom 460 colourbar. Solid, light-blue lines represent the statistical-mean position of the ultraviolet 461 auroral main oval emission⁸. For consistency with the Juno science team and the Earth-462 based community supporting the Juno mission, increasing System III longitude is shown 463 anti-clockwise. 464

Figure 2 Northern-polar projections of Jupiter's 7.80- μ m CH₄ emission. Images were recorded by Subaru-COMICS on (a-b) January 12-13, (c) February 5 and (d-f) May 18-20 2017. These are a subset of the observations shown in Supplementary Figure 1, when the northern auroral region (centred at 180°W, System III) is fully or partially visible on the disk. Images are shown in brightness temperature units according to the bottom colourbar. Solid, light-blue lines represent the statistical-mean position of the ultraviolet auroral main oval emission⁸.

Figure 3 Polar projections and regions chosen for analysis. Subaru-COMICS 7.80-μm
images recorded on (a) 2017-01-12 12:57 UT (shown in the south) and (b) 2017-01-13
12:30 UT (shown in the north), as in Figures 1 and 2, shown again here for comparison with the ultraviolet main oval statistical mean⁸ and contours that map to difference

distances in the external magnetosphere of Jupiter, as indicated in the legend. The region enclosed within the 95 R_J contour is interpreted to map to the outer magnetosphere/magnetopause. Regions A and L are respectively enclosed within the magenta and green regions and were chosen to represent the auroral and non-auroral regions for calculations of the relative radiance and its variability, as detailed in Methods.

Figure 4 Auroral-quiescent residual brightness-temperature values over time. The resid-48 ual 7.80-µm brightness temperature between Region A (the auroral region) and Region L 482 (a lower-latitude zonal mean) as described in the text/Methods are shown are red points 483 with error bars. Filled points denote results in the south, unfilled points denote results 484 in the north. Results are shown in a) January 2017 and b) May 2017. Predicted solar-485 wind dynamical pressures at Jupiter (see Methods) are shown as the solid, black line with 486 horizontal error bars showing the potential time error. The Figure suggests a brightening 487 of Jupiter's southern auroral CH₄ emission in response to a solar-wind compression at 488 approximately 22:00 UT on January 11th 2017. 489













Supplementary information to "A brightening of Jupiter's auroral 7.8- μ m CH₄ emission during a solar-wind compression" by Sinclair et al. 2019.



Supplementary Figure 1: Subaru-COMICS images of Jupiter's CH_4 emission. Images are given in chronological order from left-to-right, top-to-bottom and the mean date/time (UTC) and sub-observer System III longitude of each observation are indicated. All images are shown in 7.80-µm brightness-temperature units according to the colourbar at the bottom-right. The bright spot on the disk of Jupiter in panels h) and i) was a Galilean satellite.



Supplementary Figure 2: Mauna Kea seeing values from January 11-13th. The angular magnitude of seeing (arcsec) at 7.8 μ m on a) January 11th (UTC), b) January 12th and c) January 13th 2017. The angular magnitude was measured by the Differential Image Motion Monitor (DIMM) taken from the Mauna Kea weather center (mkwc.ifa.hawaii.edu/current/seeing/) at 0.5 μ m and scaled by a factor of (0.5/7.8)^{0.2} for wavelength dependence. Downward arrows mark the times of a subset of COMICS measurements detailed in Supplementary Table 1. These results demonstrate that atmospheric seeing was poorer on January 12th 16:13 UT compared to January 13th at 12:30 UT (~0.45"). Thus, the finer spatial structure observed in the CH₄ emission on January 12th (Figure 2a, main text) and its absence on January 13th (Figure 2b, main text) cannot be explained by poorer seeing on the latter date.



Supplementary Figure 3: The COMICS 7.80-µm bandpass. A brightness temperature spectrum of Jupiter at $\Delta \tilde{\nu} = 0.5 \text{ cm}^{-1}$ (black, solid, of the northern auroral region measured by Cassini-CIRS¹ during the 2001 flyby²), the COMICS 7.80-µm filter response (solid, red) including the telluric transmission spectrum (dotted, red) according to the red right-hand axis.



Supplementary Figure 4: Model temperature profile and functional derivatives. Panel (a) shows the vertical temperature profile (solid) and 1- σ uncertainty (dotted) retrieved from IRTF-TEXES (Texas Echelon Cross Echelle Spectrograph on NASA's Infrared Telescope Facility³) measurements of Jupiter's northern auroral region (centred at 70°W, 180°W) in December 2014⁴. Panel (b) shows the corresponding vertical functional derivatives with respect to temperature (dR_v/dT_i, where R_v is radiance at wavenumber, ν and T_i is temperature at the ith pressure level in the atmosphere). This demonstrates that the 7.8-µm bandpass is predominantly sensitive to CH₄ emission in the 20- to 0.5-mbar range with non-zero sensitivity to pressures as low as 0.1 µbar or ~360 km above the 1-bar level.



Supplementary Figure 5: Forward-modelled vertical profiles of temperature and CH₄. In panel (a), the vertical profile of CH₄ and the lapse rate in the thermosphere were both fixed while the pressure level of the mesosphere-thermosphere transition was varied as shown. In b), the vertical profile of CH₄ and the transition pressure level were fixed while the lapse rate was modified as shown. In panel (c), the vertical profile of temperature was fixed and the height of the CH₄ homopause was modified as shown. The corresponding forward-modelled brightness temperature (T_b) at 7.80 µm are indicated by the same colour in the legend of each figure. In panels (a) and (b), the temperature profile measured by Galileo during its descent⁵ is shown as the dotted, black line for comparison.

Supplementary Table 1: Details of the Subaru-COMICS images adopted in this work. All values represent the mean during the exposure time. Dates/times are given in UTC. Negative relative velocities indicate Jupiter moving towards the Earth.

Date	Time	Exposure	Angular	Airmass	Relative velocity	Filenames
(yyyy-mm-dd)	(UTC)	time (s)	diameter (")		$(\rm km/s)$	
2017-01-11	15:50	20	36.7	1.14	-28.0	wCOMA00134041
						wCOMA00134043
2017-01-12	12:57	20	36.8	1.79	-28.3	wCOMA00134551
						wCOMA00134553
	16:13	20	36.8	1.13	-28.0	wCOMA00134843
						wCOMA00134845
2017-01-13	12:30	20	36.9	2.09	-28.3	wCOMA00135473
						wCOMA00134475
2017-01-14	12:33	20	37.0	2.00	-28.3	wCOMA00136493
						wCOMA00136495
2017-02-04	14:58	20	39.4	1.13	-25.9	wCOMA00138283
						wCOMA00138285
2017-02-05	15:54	20	39.6	1.18	-25.6	wCOMA00139653
						wCOMA00139655
2017-05-17	09:02	20	42.2	1.18	18.19	wCOMA00139989
						wCOMA00139991
						wCOMA00139993
						wCOMA00139995
2017-05-17	09:40	20	42.2	1.29	18.3	wCOMA00140085
						wCOMA00140087
						wCOMA00140089
						wCOMA00140091
2017-05-18	09:35	20	42.1	1.28	18.3	wCOMA00141011
						wCOMA00141013
						wCOMA00141015
						wCOMA00141017
2017-05-19	05:37	20	42.1	1.24	18.3	wCOMA00141385
						wCOMA00141387
						wCOMA00141389
						wCOMA00141391
2017-05-20	05:54	20	42.0	1.18	18.3	wCOMA00142085
						wCOMA00142087
						wCOMA00142089
						wCOMA00142091
2017-05-20	06:30	20	42.0	1.12	18.3	wCOMA00142181
						wCOMA00142183
						wCOMA00142185
						wCOMA00142187
2017-05-20	09:55	20	41.90	1.40	18.3	wCOMA00142675
						wCOMA00142677
						wCOMA00142679
						wCOMA00142681
2017-05-20 2017-05-20	06:30 09:55	20	42.0 41.90	1.12	18.3	wCOMA00142181 wCOMA00142183 wCOMA00142185 wCOMA00142187 wCOMA00142675 wCOMA00142677 wCOMA00142679 wCOMA00142681

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